

PRELIMINARY NOTE ON THE LITHOGEOCHEMISTRY AND PETROGENESIS OF INTRUSIVE ROCK SUITES FROM THE MINIPI LAKE REGION (NTS MAP AREA 13C/SOUTH), GRENVILLE PROVINCE, LABRADOR

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ABSTRACT

The Minipi Lake region of southern Labrador straddles the boundary of two major Grenvillian lithotectonic divisions: the Mealy Mountains terrane and the Mecatina terrane. Both are dominated by variably deformed and metamorphosed igneous suites containing local screens of older ortho- and paragneisses.

Four contrasting intrusive rock groups have been distinguished, including: 1) the late Paleoproterozoic Mealy Mountains intrusive suite (MMIS), comprising rock types varying from gabbro to pyroxene monzonite, and occurring in the Mealy Mountains terrane, 2) early Mesoproterozoic units of variably foliated quartz monzonite and K-feldspar porphyritic granite, which underlie much of the Mecatina terrane in the study area, 3) a distinct group of gabbro and gabbro-norite, forming small plugs and dykes that punctuate and appear to be restricted to the Mecatina terrane, and 4) plutons of late- to post-Grenvillian granite and K-feldspar porphyritic granite that occur in the Mealy Mountains and Mecatina terranes.

Samples of the MMIS display remarkably similar geochemical signatures. Except for one, all are olivine normative, most are silica-saturated (olivine + hypersthene in the norm) and only a few can be qualified as silica-undersaturated (olivine + nepheline in the norm). As a group, these rocks can be described as transitional to alkaline. Ultramafic and mafic samples are markedly enriched in Al_2O_3 and deficient in TiO_2 compared to basalt and diabase.

Trace-element patterns in MMIS rocks show a marked enrichment in high-field-strength elements (HFSE) relative to N-MORB, and a relatively flat heavy-REE distribution near, or somewhat below the normalizing values. The most notable features of the patterns are given by the well-defined negative Th, Nb, Zr, Hf, and Ti anomalies, which are coupled with a positive Sr spike. The same trace-element pattern is duplicated in all the analyses, independently of the silica content of the rock, hence providing strong support for a cogenetic origin. Accordingly, the compositional spectrum of the suite can be attributed to variable degrees of fractionation. The patterns also emphasize the importance that fractional crystallization and extraction of cumulates could have played in the petrogenesis of the suite.

The geochemical signature of the Mecatina terrane gabbro and gabbro-norite differs significantly from that of the MMIS. Their major-element contents are comparable to those of basalt and diabase, except for a small enrichment in Al_2O_3 and small depletions in TiO_2 and P_2O_5 that are consistent with limited plagioclase fractionation. However, incompatible trace-elements Th, Zr, Hf and Y show no evidence of fractionation. These samples plot within the volcanic-arc basalt field in Ti-Zr-Y and Hf-Th-Ta tectonic discrimination diagrams.

The early Mesoproterozoic quartz monzonite and K-feldspar porphyritic granite samples from the Mecatina terrane plot in the volcanic-arc granite field on Nb-Y and Rb-(Nb+Y) discrimination diagrams. Taken together, the volcanic-arc signatures of Mecatina terrane gabbroic, and spatially associated, granitic rocks provide support for models involving the development of an Andean-type margin in early Mesoproterozoic. In contrast, the within-plate granite signature of the late- to post-Grenvillian granites is consistent with their emplacement in significantly older crust.

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INTRODUCTION AND REGIONAL GEOLOGICAL SETTING

The Minipi Lake region (NTS map area 13C/South) of southern Labrador includes parts of two major Grenvillian lithotectonic divisions, the Mealy Mountains terrane and the Mecatina terrane (Figures 1 and 2). The geology and field relationships of the different plutonic rock suites, which underlie most of the two terranes, have recently been described by James and Nadeau (2000a, b). Preliminary U–Pb geochronological data are reported by James *et al.* (*this volume*). Focus of this complementary note is placed on the whole-rock, major- and trace-element geochemical results obtained from a set of representative rock samples collected during mapping.

The northern two-thirds of the region (Figure 1), part of the Mealy Mountains terrane (Gower and Owen, 1984), is dominated by massive to weakly deformed and recrystallized gabbroic and monzodioritic rocks of the late Paleoproterozoic (Labradorian-age) Mealy Mountains intrusive suite (MMIS) (*see* Emslie, 1976; and Emslie and Hunt, 1990). In contrast, the area south of the Little Mecatina River (Figure 1), provisionally correlated with the northern extension of the Mecatina terrane (Figure 2), is mainly underlain by early Mesoproterozoic (Pinwarian-age), variably foliated and commonly porphyritic, granite and quartz monzonite that are intruded by small plugs and dykes of gabbro and gabbro-norite. The Mealy Mountains and Mecatina terranes are intruded by plutons consisting of mostly massive, medium- to coarse-grained granite and K-feldspar porphyritic granite of late- to post-Grenvillian age.

SAMPLING AND ANALYTICAL METHODS

Fifty-six samples, representative of all major rock types, were selected for major- and trace-element analysis on the basis of their mesoscopic textural homogeneity and representativeness of the outcrop area (Figure 2; Tables 1 and 2). None of the samples show field evidence of metamorphic remobilization (e.g., migmatization), fracture related alteration, or weathering. Textures of the rock samples are varied from massive, having complete preservation of igneous texture and mineralogy, to metamorphically recrystallized and foliated. Most gabbroic samples are medium to coarse grained, whereas granitic samples are varied from coarse grained to K-feldspar porphyritic. With a few exceptions, each sample consisted of more than two kilograms of fresh rock.

Major elements, transition metals (Sc, V, Cr, Co, Ni), and Cu, Zn, Sr and Ba were determined by Inductively Coupled Plasma - Atomic Emission Spectrometry (ICP-AES).

Rare-earth elements (REE), Pb, Y, Hf, Zr, Rb, Ta, and Th were analyzed by Inductively Coupled Plasma - Mass Spectrometry (ICP-MS). Detection limits are given in Tables 1 and 2. Samples were low in volatiles with average ($\text{H}_2\text{O}+\text{CO}_2$) of 0.76 wt.%, and most totals between 98.7 wt.% and 101.3 wt.%. Norm (CIPW) have been calculated volatile-free to 100% (Table 3). Total iron has been expressed as FeO^* using the conversion factor of Irvine and Baragar (1971). Analyses were performed at the geochemistry laboratory of the Québec Geoscience Centre.

MAJOR INTRUSIVE SUITES

MEALY MOUNTAINS INTRUSIVE SUITE

The MMIS consists of an older group of anorthositic, leucogabbroic and leucotroctolitic rocks, and a younger group of pyroxene-bearing monzodioritic to granitic rocks. It constitutes one of the most extensive intrusive suites of the Grenville Province in Labrador, underlying most of the region extending from the highlands southeast of Lake Melville, through the Minipi Lake region, over more than 200 km in the southwest direction (*cf.* Emslie, 1976; Gower and van Nostrand, 1996; Nunn and van Nostrand, 1996; Gower, 1999; and James and Lawlor, 1999). In the Minipi Lake region (Figure 1), the MMIS includes two, regionally extensive rock units consisting mainly of gabbro and gabbro-norite (Unit P_{MM} gbr), and of pyroxene-bearing monzodiorite to monzonite (Unit P_{MM} mdq).

As described by James and Nadeau (2000a), MMIS gabbro and gabbro-norite (Unit P_{MM} gbr, Figure 1), have a wide range in compositions, textures, structures and aeromagnetic signatures, possibly indicating that the unit consists of intrusions of different ages. The unit also contains subordinate amounts of leucogabbro, leucogabbro-norite, pyroxenite, diorite and very minor amounts of monzogabbro, monzogabbro-norite and amphibolite; anorthosite nor leucotroctolite occur in the study area. Rocks consist of varied amounts of clinopyroxene, orthopyroxene and plagioclase; biotite is a common accessory mineral. The rocks are varied, from massive and unrecrystallized (containing preserved igneous mineral textures) to foliated and pervasively recrystallized. Trace amounts to several percent of pyrite are common.

James and Nadeau (2000a) describe the MMIS pyroxene-bearing monzodiorite and monzonite unit (Unit P_{MM} mdq) as a compositionally and texturally varied unit including monzodiorite, quartz monzodiorite, diorite, granodiorite quartz monzonite and granite. The unit also includes minor amounts of gabbro and gabbro-norite that are correlated with Unit P_{MM} gbr. Rocks are texturally varied, even at the outcrop scale, from massive with complete preservation of

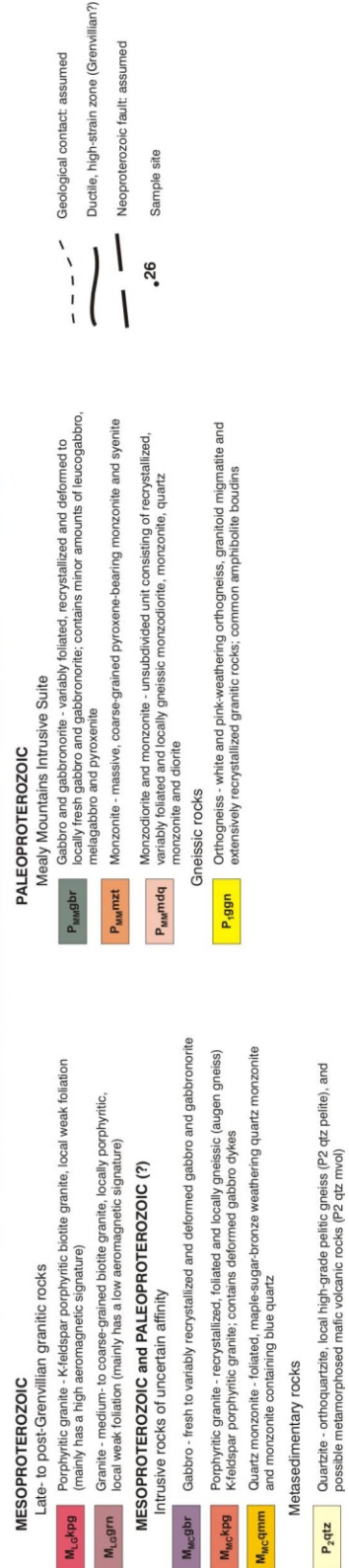


Figure 1. Geological map of the Mirapi Lake region, NTS map area 13C/South (including NTS map areas 13C/1, C/2, C/3, C/6, C/7 and C/8). Simplified from James and Nadeau (2000b).

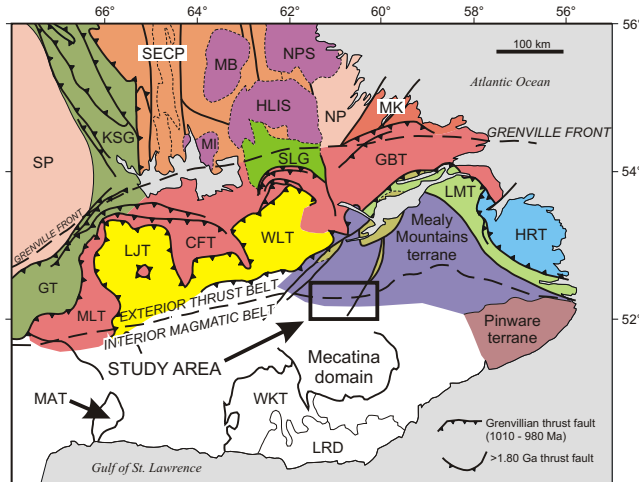


Figure 2. Location of the Minipi Lake region in relation to the tectonic and major lithotectonic divisions of northeastern Laurentia (modified from Wardle *et al.*, 1997). Grenville Province: HRT - Hawke River terrane, LMT - Lake Melville terrane, GBT - Groswater Bay terrane, WLT - Wilson Lake terrane, CFT - Churchill Falls terrane, LJT - Lac Joseph terrane, MLT - Molson Lake terrane, GT - Gagnon terrane, MAT - Matamec terrane, WKT - Wakeham terrane, LRD - La Romaine terrane. Archean divisions: SP - Superior Province, NP - Nain Province (Hopedale Block). Archean and Paleoproterozoic divisions: MK - Makkovik Province, SECP - Southeastern Churchill Province (Core Zone), KSG - Kaniapiskau Supergroup (2.25-1.86 Ga). Mesoproterozoic units: NPS - Nain Plutonic Suite, HLIS - Harp Lake intrusive suite, MB - Mistastin batholith, MI - Michikamau Intrusion, SLG - Seal Lake Group.

igneous textures, to recrystallized and foliated; gneissic varieties also occur. The unit probably consists of distinct intrusions of relatively uniform composition, although it is so poorly exposed that it cannot be subdivided at the present scale of mapping. Clinopyroxene, hornblende, biotite and accessory magnetite are common in these rocks. A sample of gneissic P_{MM} mdq monzodiorite, which occurs near Minipi Lake (NTS map area 13C/10), has an igneous emplacement age of 1659 ± 5 Ma based on U-Pb dating of zircon (James *et al.*, 2000).

MECATINA TERRANE QUARTZ MONZONITE AND GRANITE

Quartz monzonite (Unit M_{MC} qmm) and K-feldspar porphyritic granite (Unit M_{MC} kpg) underlie much of Mecatina terrane in the study area. Samples from these units have yielded U-Pb zircon igneous crystallization ages of ca. 1500 ± 4 Ma and 1493 ± 3 Ma, respectively (James *et al.*, *this volume*). (The quartz monzonite sample 43 and porphyritic granite sample 46 (see Figure 1) are from the same

outcrops that were sampled for the geochronological study.) Unit M_{MC} qmm consists of variably foliated, medium- to coarse-grained, equigranular to K-feldspar porphyritic quartz monzonite and, less commonly, monzonite. Quartz is commonly pale blue. These rocks contain less than 10 percent mafic minerals including clinopyroxene, biotite and accessory magnetite. The K-feldspar porphyritic granite (Unit M_{MC} kpg) is generally well foliated and contains less than 10 percent combined biotite and hornblende.

MECATINA TERRANE GABBRO

Two small bodies and dykes of gabbro (Unit M_{MC} gbr) intrude the quartz monzonite and K-feldspar porphyritic granite units of the Mecatina terrane (Figure 1). The unit mostly consists of fresh gabbro and gabbronorite, although rocks are locally recrystallized and foliated. The rocks are typically medium to coarse grained, ophitic to subophitic. They are mainly homogeneous in composition, although they locally display remarkable igneous layering (James and Nadeau, 2000a). These rocks are interpreted to be part of the regionally extensive Petit Mecatina anorthosite-mangerite-charnockite-granite (AMCG) complex underlying much of the Mecatina terrane south of the study area. Three of the five samples analyzed come from small dykes contained in quartz monzonite and K-feldspar porphyritic granite.

LATE- TO POST-GRENVILLIAN GRANITE

The Mealy Mountains and Mecatina terranes are intruded by plutons consisting of medium- to coarse-grained granite (Unit M_{LG} grn) and K-feldspar porphyritic granite (Unit M_{LG} kpg); the two units have similar composition. Rocks are mainly undeformed and massive, and contain less than 10 percent biotite. Similar granite plutons punctuate the Mealy Mountains and the Pinware terranes east of the study area and have emplacement ages between 980 and 950 Ma (see Gower, 1996; Gower *et al.*, 1991). A sample of coarse-grained Unit M_{LG} grn granite occurring in the northeastern part of the Minipi Lake region has an emplacement age of 964 ± 3 Ma based on U-Pb dating of zircon (James *et al.*, *this volume*).

RESULTS AND DISCUSSION

It is emphasized that most gabbroic to monzodioritic rocks exhibit medium- and coarse-grained igneous or relict igneous textures, locally with a distinct igneous mineral foliation and, more rarely, compositional layering. Granitic rocks are generally coarse grained, equigranular or K-feldspar porphyritic, and locally retain a subtle prismatic mineral foliation, possibly of igneous origin. These primary igneous textural attributes indicate that some crystal separation and settling occurred during crystallization. Therefore,

Table 2. Trace-element contents of intrusive rock samples from the Mirapi Lake region (For P_{MMBgr} read P_{MMgbr}.)

Site	Sample number	Element Method Limit	Rb	Sr	Ba	Nb	Zr	Y	Th	La	Ce	Hf	Nd	Sm	Eu	Tb	Yb	Gd	Dy	Ho	Er	Tm	Yb	Lu			
			ICP-MS	ICP-AES	ICP-AES	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS	ICP-MS			
			ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm			
Mealy Mountains intrusive suite																											
24	DJ-99-059	P _{MMBgr}	27	1256	1999	9	90	29	0.55	43.26	105.44	2.09	73.59	14.76	4.54	1.34	1.58	12.36	6.71	1.10	2.61	0.30	0.54	15.24	0.24		
25	NK-99-056C	P _{MMBgr}	29	739	1365	3	84	57	0.38	51.88	121.66	1.99	81.31	17.17	4.62	1.89	3.96	15.35	10.81	2.03	5.47	0.70	1.15	17.31	0.59		
11	DJ-99-187	P _{MMBgr}	29	1385	1210	3	30	24	0.09	13.27	36.85	1.08	28.15	6.47	2.14	0.78	2.14	5.94	4.67	0.91	2.55	0.36	0.09	5.80	0.33		
28	NK-99-040	P _{MMBgr}	11	1174	352	1	22	21	0.29	10.83	28.02	0.69	21.90	5.28	1.43	0.63	1.66	4.87	3.76	0.72	1.98	0.27	0.04	4.39	0.25		
3	NK-99-068B	P _{MMBgr}	19	1499	792	2	32	15	0.27	16.75	38.40	0.89	23.73	4.88	1.61	0.53	1.25	4.31	3.05	0.58	1.62	0.21	0.09	5.26	0.19		
22	NK-99-050	P _{MMBgr}	8	1394	1156	9	109	22	0.32	45.25	102.12	2.30	59.76	10.94	3.36	0.94	1.23	8.71	4.69	0.79	1.92	0.23	0.46	13.61	0.18		
13	NK-99-053B	P _{MMBgr}	12	1815	1214	4	40	24	0.12	19.13	48.53	1.15	33.09	7.03	2.15	0.75	1.77	5.92	4.28	0.81	2.21	0.31	0.11	7.19	0.28		
29	NK-99-038	P _{MMBgr}	10	506	252	1	10	9	0.10	3.24	7.28	0.31	5.47	1.46	0.76	0.23	0.67	1.66	1.50	0.31	0.83	0.11	0.03	1.07	0.10		
18	DJ-99-176	P _{MMBgr}	15	393	324	8	124	29	0.93	11.85	28.27	3.05	18.72	4.84	1.71	0.81	2.66	5.29	5.22	1.05	3.06	0.44	0.55	3.99	0.40		
40	DJ-99-141	P _{MMBgr}	17	320	244	12	173	39	1.26	12.50	31.78	4.04	21.79	6.07	2.07	1.08	3.56	6.96	7.00	1.42	4.13	0.59	0.80	4.59	0.56		
20	NK-99-051B	P _{MMBgr}	25	358	606	10	191	39	1.62	19.92	46.91	4.58	29.31	7.21	2.22	1.14	3.43	7.58	7.08	1.43	4.06	0.57	0.62	6.47	0.53		
7	DJ-99-191	P _{MMBgr}	8	292	232	6	89	23	0.57	7.96	19.47	2.21	13.23	3.63	1.35	0.65	2.14	4.13	4.24	0.85	2.44	0.35	0.42	2.80	0.33		
26	DJ-99-253	P _{MMBgr}	1	552	197	0	7	7	0.05	1.80	4.51	0.24	3.89	1.14	0.60	0.20	0.59	1.30	1.32	0.26	0.74	0.10	0.01	0.74	0.09		
21	NK-99-049	P _{MMBgr}	5	1304	417	2	27	13	0.12	7.43	18.07	0.79	13.08	3.26	0.97	0.39	1.09	2.85	2.40	0.46	1.33	0.19	0.06	2.70	0.17		
37	DJ-99-154	P _{MMBgr}	11	421	456	7	125	26	0.85	13.38	31.22	2.95	19.99	4.96	1.73	0.76	2.29	5.26	4.81	0.95	2.71	0.39	0.41	4.34	0.35		
16	NK-99-054A	P _{MMBgr}	21	1311	648	1	25	9	0.53	8.65	17.48	0.61	9.93	2.16	0.93	0.27	0.73	2.02	1.59	0.30	0.90	0.12	0.05	2.29	0.12		
15	NK-99-061B	P _{MMBgr}	27	2781	799	4	30	11	1.23	21.36	45.47	0.68	22.56	3.97	1.30	0.36	0.86	3.00	2.00	0.37	1.03	0.14	0.15	5.70	0.14		
23	NK-99-045B	P _{MMBgr}	27	1684	2624	4	43	46	0.64	48.81	108.06	1.06	62.78	12.05	3.69	1.33	3.40	10.58	7.84	1.55	4.35	0.58	0.18	14.43	0.52		
22	NK-99-046B	P _{MMBgr}	22	1688	1263	2	14	17	0.35	24.09	50.86	0.36	27.40	5.16	1.78	0.53	1.25	4.31	2.93	0.56	1.54	0.21	0.08	6.52	0.20		
34	NK-99-034B	P _{MMBgr}	7	2146	545	1	17	8	0.24	8.04	16.69	0.40	10.05	2.19	0.95	0.25	0.66	1.89	1.44	0.28	0.78	0.11	0.04	2.23	0.10		
8	NK-99-118	P _{MMBgr}	10	2000	883	1	22	14	0.62	19.19	42.93	0.63	25.56	5.15	1.66	0.51	1.12	4.21	2.84	0.53	1.47	0.19	0.06	5.77	0.18		
1	NK-99-135B	P _{MMBgr}	18	1330	1126	2	32	20	0.80	26.42	58.01	0.91	32.63	6.56	1.81	0.66	1.77	5.49	3.82	0.74	2.08	0.29	0.11	7.51	0.28		
2	DJ-99-281B	P _{MMBgr}	10	1704	1162	2	23	16	0.65	24.54	53.17	0.65	29.40	5.61	1.88	0.56	1.33	4.67	3.16	0.59	1.64	0.23	0.09	6.90	0.21		
35	NK-99-033B	P _{MMBgr}	15	1340	1008	3	26	24	0.37	31.02	65.02	0.65	34.26	6.59	1.91	0.69	1.92	5.54	4.00	0.79	2.21	0.31	0.11	8.28	0.29		
30	NK-99-039	P _{MMBgr}	56	1374	4899	3	17	14	0.53	27.70	54.18	0.40	27.46	5.00	2.62	0.49	0.95	4.11	2.65	0.51	1.35	0.18	0.13	6.66	0.15		
36	DJ-99-155	P _{MMBgr}	13	1585	1534	1	11	13	0.41	27.42	52.79	0.29	26.37	4.73	2.29	0.46	0.98	3.88	2.52	0.47	1.29	0.17	0.06	6.37	0.15		
33	DJ-99-151	P _{MMBgr}	45	1115	1531	5	223	17	1.49	26.59	56.62	4.54	29.84	5.57	1.69	0.55	1.60	4.58	3.06	0.60	1.75	0.26	0.21	7.13	0.27		
9	NK-99-133	P _{MMBgr}	19	1976	1252	2	25	17	1.00	29.43	62.21	0.69	31.80	5.80	2.02	0.56	1.40	4.71	3.17	0.61	1.73	0.23	0.10	7.76	0.21		
17	NK-99-055B	P _{MMBgr}	52	1859	1767	6	68	17	1.14	34.57	70.33	1.48	34.96	6.22	2.13	0.58	1.37	4.83	3.14	0.59	1.66	0.23	0.23	8.68	0.22		
19	NK-99-044B	P _{MMBgr}	34	891	1561	6	139	28	0.42	31.58	65.91	2.66	34.02	6.74	1.75	0.76	2.41	5.72	4.51	0.91	2.64	0.39	0.20	8.28	0.38		
10	DJ-99-279B	P _{MMBgr}	17	1377	1778	2	23	17	0.71	25.33	53.17	0.70	28.38	5.61	1.89	0.56	1.53	4.53	3.22	0.62	1.73	0.25	0.10	6.75	0.25		
4	NK-99-057B	P _{MMBgr}	52	1481	6895	4	109	14	0.43	30.77	60.85	2.20	29.45	5.12	2.59	0.49	1.16	4.08	2.76	0.52	1.49	0.20	0.15	7.31	0.18		
14	NK-99-063B	P _{MMBgr}	53	821	1938	9	70	24	1.76	32.34	65.35	1.74	33.98	6.60	2.11	0.71	2.18	5.49	4.18	0.83	2.46	0.35	0.35	8.28	0.35		
5	NK-99-066B	P _{MMBgr}	68	990	1716	6	115	22	4.46	37.44	77.00	3.21	37.60	7.02	1.71	0.70	2.24	5.49	4.24	0.82	2.42	0.35	0.39	9.55	0.36		
Petit Mecatina domain intrusive suite																											
48	DJ-99-098	M _{Mc} gbr	152	229	152	4	57	18	1.16	5.89	15.11	1.42	9.50	2.58	1.08	0.48	1.63	3.05	3.12	0.65	1.86	0.27	0.21	6.03	0.26		
47	NK-99-020B	M _{Mc} gbr	32	314	443	11	203	39	2.06	21.97	49.85	4.63	29.71	7.12	2.04	1.09	3.42	7.36	6.94	1.39	3.98	0.56	0.66	6.78	0.52		
41	NK-99-029B	M _{Mc} gbr	42	337	533	14	252	45	2.73	28.01	63.06	5.59	36.27	8.43	2.34	1.26	3.85	8.65	7.85	1.58	4.48	0.64	0.82	8.37	0.59		
49	NK-99-009B	M _{Mc} gbr	20	292	341	9	153	34	1.60	16.40	37.01	3.80	23.00	5.81	1.84	0.94	3.11	6.34	6.21	1.24	3.63	0.51	0.51	5.10	0.48		
50	NK-99-012B	M _{Mc} gbr	2	330	41	3	54	16	0.43	4.07	9.54	1.40	6.76	2.02	0.74	0.40	1.40	2.58	2.74	0.57	1.71	0.24	0.17	1.39	0.23		
43	DJ-99-084	M _{Mc} qmm	51	521	2545	16	458	55	1.96	47.94	103.42	9.12	57.34	11.96	3.38	1.57	4.76	11.28	9.62	1.94	5.57	0.78	0.73	13.43	0.75		
42	NK-99-021B	M _{Mc} qmm	100	641	1855	7	202	20	1.73	37.27	72.79	4.38	33.81	6.20	1.74	0.61	1.84	4.84	3.59	0.70	2.02	0.30	0.29	8.67	0.30		
51	NK-99-005B	M _{Mc} kpg	151	272	1773	14	412	52	5.43	34.19	77.77	9.52	43.21	9.38	2.30	1.32	5.43	8.90	8.85	1.80	5.69	0.86	0.95	10.36	0.86		
46	DJ-99-110	M _{Mc} kpg	208	144	1016	12	178	19	37.19	47.58	101.48	4.69	31.27	4.55	0.81	0.51	2.43	3.76	3.32	0.68	2.12	0.33	0.88	9.71	0.58		
45	DJ-99-131	M _{Mc} kpg	21																								

Table 3. Table of the CIPW norm of intrusive rock samples from the Minipi Lake region

Site	Sample number	Lithology	Normative mineral	Q	or	ab	an	PLAG	n	Di wo	Di en	Difs	CPX	Hy en	Hy fs	OPX	Oi fo	Oi fa	OL	mt	il	ap
Mealy Mountains intrusive suite																						
24	DJ-99-059	monzogabbro		23.01	6.15	20.65	43.66	2.22	9.61	2.97	2.62	5.59	7.13	6.94	14.07	4.86	8.70	7.35	8.70	4.86	8.70	7.35
25	NK-99-056C	gabbro		20.59	8.43	17.90	38.49	3.87	13.30	6.64	2.79	3.87	3.87	14.03	23.21	5.00	9.18	14.03	23.21	5.00	9.18	14.03
11	DJ-99-187	gabbro		6.35	10.97	37.78	44.13	2.13	8.56	7.92	2.11	2.13	8.56	10.07	21.26	3.47	10.07	21.26	3.47	3.47	2.42	1.20
28	NK-99-040	gabbro		3.62	17.86	30.58	48.44	3.92	15.76	4.35	3.89	3.92	15.76	2.14	20.28	3.98	9.60	10.68	20.28	3.98	2.29	1.38
3	NK-99-068B	dioritic gneiss		7.57	20.21	34.31	54.52	1.81	6.74	3.39	1.54	1.81	6.74	9.00	20.67	3.52	9.00	11.67	20.67	3.52	2.37	1.32
22	NK-99-050	gabbro		30.67	6.75	25.67	56.34	2.02	7.95	4.02	2.02	1.91	7.95	0.22	16.43	3.70	6.85	6.85	16.43	3.70	6.33	5.05
13	NK-99-053B	monzodioritic/dioritic gneiss		20.45	6.85	34.56	55.01	1.75	6.54	3.28	1.51	1.75	6.54	8.62	19.64	3.28	8.62	11.02	19.64	3.28	2.10	0.92
29	NK-99-038	gabbro		16.34	3.56	35.84	52.18	5.13	14.77	7.68	5.13	1.96	14.77	4.14	19.95	2.17	14.04	5.91	19.95	2.17	1.08	0.58
18	DJ-99-176	gabbro		23.17	4.98	31.22	54.39	2.13	9.50	4.82	2.55	2.13	9.50	3.70	16.79	3.60	8.74	8.05	16.79	3.60	3.39	0.56
40	DJ-99-141	gabbro		29.20	4.23	29.03	58.23	2.46	10.48	5.27	2.46	2.75	10.48	0.65	16.64	3.88	7.45	9.19	16.64	3.88	4.53	0.65
7	DJ-99-151	gabbro/clino pyroxenite		23.15	7.89	24.65	47.80	2.73	10.81	5.47	2.73	2.61	10.81	4.65	14.62	4.19	7.11	7.51	14.62	4.19	4.74	0.87
20	NK-99-051B	pyroxenite		21.02	3.23	32.09	53.11	3.75	13.19	6.73	3.75	2.71	13.19	3.25	18.06	3.43	10.12	8.06	18.06	3.43	2.90	0.38
26	DJ-99-253	gabbro		11.61	0.83	36.34	47.95	2.69	21.51	11.20	7.62	2.69	21.51	1.67	24.80	2.18	17.85	6.95	24.80	2.18	0.48	-
21	NK-99-049	gabbro		14.87	3.33	44.24	59.11	2.96	19.91	10.31	6.64	2.96	19.91	1.78	12.24	1.75	8.20	4.04	12.24	1.75	0.89	0.20
37	DJ-99-154	gabbro		23.35	4.44	30.06	53.41	3.06	10.77	5.90	3.06	2.21	10.77	4.29	16.62	3.44	9.25	7.37	16.62	3.44	3.31	0.61
16	NK-99-054A	gabbro		19.35	5.75	41.34	60.69	2.00	12.35	6.37	3.98	2.00	12.35	3.51	12.99	1.85	4.63	4.63	12.99	1.85	0.79	0.31
15	NK-99-061B	dioritic gneiss		28.08	7.65	50.64	78.72	0.31	6.74	1.40	0.65	0.74	6.74	1.76	7.19	1.22	3.19	4.00	7.19	1.22	0.80	0.32
23	NK-99-045B	monzonite		36.64	13.61	20.16	56.80	0.49	1.67	0.83	0.35	0.49	1.67	1.35	14.32	3.11	5.61	8.71	14.32	3.11	2.89	4.35
12	NK-99-046B	monzonite		34.36	7.18	31.13	65.49	1.15	4.47	2.25	1.07	1.15	4.47	1.88	15.65	2.50	7.15	8.50	15.65	2.50	1.67	1.23
34	NK-99-034B	monzogabbro/monzodiorite		30.18	4.75	39.45	69.63	1.43	6.09	3.06	1.43	1.60	6.09	1.69	12.63	2.29	5.65	6.98	12.63	2.29	0.72	0.33
8	NK-99-118	diorite/gabbro		29.17	4.93	34.30	63.47	0.21	11.95	6.08	3.30	2.57	11.95	7.87	14.63	2.24	7.87	6.76	14.63	2.24	1.22	1.36
1	NK-99-135B	gabbro		28.43	7.74	25.79	54.22	4.06	14.73	7.50	4.06	3.17	14.73	4.46	10.09	2.56	5.42	5.67	10.09	2.56	1.55	1.19
2	DJ-99-281B	biotite gabbro		32.61	7.06	29.53	62.14	3.00	11.79	5.96	3.00	2.83	11.79	1.50	10.98	2.33	5.38	3.19	10.98	2.33	1.45	1.35
35	NK-99-033B	gabbro		36.32	7.24	29.34	65.66	1.66	6.74	3.40	1.68	1.66	6.74	0.50	6.12	2.19	4.80	4.73	6.12	2.19	1.56	0.97
30	NK-99-039	monzogabbro/monzodiorite		40.75	18.03	27.00	67.75	-	-	-	-	-	-	0.51	7.92	1.46	3.75	4.17	7.92	1.46	1.64	0.91
36	DJ-99-155	leucogabbro		42.51	9.25	32.96	75.47	0.54	3.71	1.85	0.77	1.09	3.71	2.83	7.47	1.46	2.83	4.54	7.47	1.46	1.05	1.05
33	DJ-99-151	gabbro		37.99	15.22	22.99	60.98	0.99	3.78	1.90	0.89	0.99	3.78	2.17	10.53	2.25	4.73	5.80	10.53	2.25	1.51	1.15
9	NK-99-133	meta-gabbro		39.87	8.67	31.33	71.20	0.93	3.60	1.81	0.86	0.93	3.60	2.88	6.15	1.74	6.02	2.79	6.15	1.74	1.18	1.42
17	NK-99-055B	meta-gabbro		39.80	13.62	27.40	67.20	1.19	5.00	2.53	1.28	1.19	5.00	4.62	9.37	1.61	4.62	4.75	9.37	1.61	1.34	0.99
19	NK-99-044B	monzogabbro/monzodiorite		37.42	14.34	22.90	60.32	1.22	5.00	2.53	1.25	1.22	5.00	8.12	16.05	-	0.55	0.44	16.05	-	1.91	1.39
10	DJ-99-279B	biotite gabbro		34.77	11.87	23.92	58.69	0.27	6.74	4.67	2.61	1.88	9.16	9.01	15.50	1.79	3.25	4.29	15.50	1.79	1.08	0.93
4	NK-99-067B	qtz-monzonitic gneiss		40.24	19.87	26.48	67.42	0.27	6.74	2.20	1.01	1.17	4.38	4.19	7.54	1.33	3.25	4.29	7.54	1.33	1.39	0.84
14	NK-99-063B	monzogabbro/monzodiorite		40.27	19.84	60.11	-	-	-	-	-	-	-	-	2.73	1.70	1.20	1.53	2.73	1.70	1.33	0.84
5	NK-99-066B	monzodioritic/dioritic gneiss		36.02	18.73	54.75	-	-	-	1.12	0.47	0.66	2.25	5.45	13.11	-	-	-	13.11	-	1.70	1.14
Petit Mecatina domain intrusive suite																						
48	DJ-99-098	gabbro		18.43	8.62	33.74	52.17	2.15	12.94	6.67	4.12	2.15	12.94	0.87	20.05	2.61	12.71	7.34	20.05	2.61	2.08	0.20
47	NK-99-020B	gabbro		25.41	6.91	26.83	52.24	2.77	11.95	6.06	3.12	2.77	11.95	3.96	12.61	3.68	6.37	6.24	12.61	3.68	4.40	0.73
41	NK-99-029B	gabbro		26.20	9.26	25.78	51.98	2.54	10.12	5.10	2.48	2.54	10.12	3.76	11.48	3.79	5.39	6.09	11.48	3.79	4.83	0.93
49	NK-99-009B	gabbro		20.87	4.87	28.95	49.82	3.02	14.17	7.22	3.93	3.02	14.17	11.33	20.05	2.10	1.78	3.88	20.05	2.10	3.32	3.38
50	NK-99-012B	gabbro		20.98	1.06	36.22	57.20	2.11	14.50	7.51	4.88	2.11	14.50	10.65	15.26	2.18	5.29	2.53	15.26	2.18	1.78	0.20
43	DJ-99-084	qtz-monzonite		36.26	17.36	16.08	52.34	0.80	2.54	1.26	0.48	0.80	2.54	4.77	7.93	2.21	8.12	7.93	7.93	2.21	2.67	1.16
42	NK-99-021B	qtz-monzonite gneiss		36.62	10.92	14.09	50.71	0.28	3.4	0.62	0.28	0.34	1.24	3.69	4.44	1.05	0.44	8.13	4.44	1.05	0.94	0.47
51	NK-99-005B	granitic gneiss		33.11	20.30	7.52	40.63	0.54	1.61	0.80	0.27	0.54	1.61	2.06	6.15	1.15	3.25	4.29	6.15	1.15	1.36	0.54
46	DJ-99-110	ksp-phy. Granite		35.85	28.79	26.29	3.64	29.93	0.55	0.55	0.16	0.41	1.12	0.85	2.11	2.96	-	-	2.11	2.96	0.60	0.13
45	DJ-99-131	granite		26.80	40.63	1.55	28.35	0.26	0.56	0.27	0.03	0.26	0.56	0.15	1.16	1.31	-	-	1.16	1.31	0.30	0.02
Late- to post- Grenvillian granite																						
39	NK-99-078	porphyritic granite		34.53	9.10	8.01	42.54	0.96	3.02	1.50	0.56	0.96	3.02	2.97	5.11	8.08	-	-	5.11	8.08	1.64	1.27
27	NK-99-037	biotite augen granite		28.91	34.01	6.04	34.95	0.42	1.29	0.64	0.23	0.42	1.29	1.84	3.33	5.17	-	-	3.33	5.17	1.05	0.71
32	NK-99-072B	biotite granite		32.01	15.17	8.97	40.98	0.80	2.41	1.19	0.42	0.80	2.41	2.39	4.55	6.94	-	-	4.55	6.94	1.44	1.67
Plutonic rock of uncertain affinity																						
6	NK-99-064B	biotite augen granite		37.28	13.75	8.85	46.13	-	-	-	-	-	-	2.87	3.49	6.36	-	-	3.49	6.36	0.98	0.58
38	NK-99-126B	biotite granite		42.33	34.90	7.07	49.40	-	-	-	-	-	-	0.82	1.55	2.37	-	-	1.55	2.37	0.36	0.16
53	DJ-99-091	foliated granite		29.28	31.38	32.65	29.28	1.71	30.99	0.65	0.10	0.61	3.36	0.35	2.07	2.42	-	-	2.07	2.42	0.59	0.50
16	NK-99-054F	felsic dyke		45.82	3.87	45.82	21.98	67.60	-	0.27	0.12											

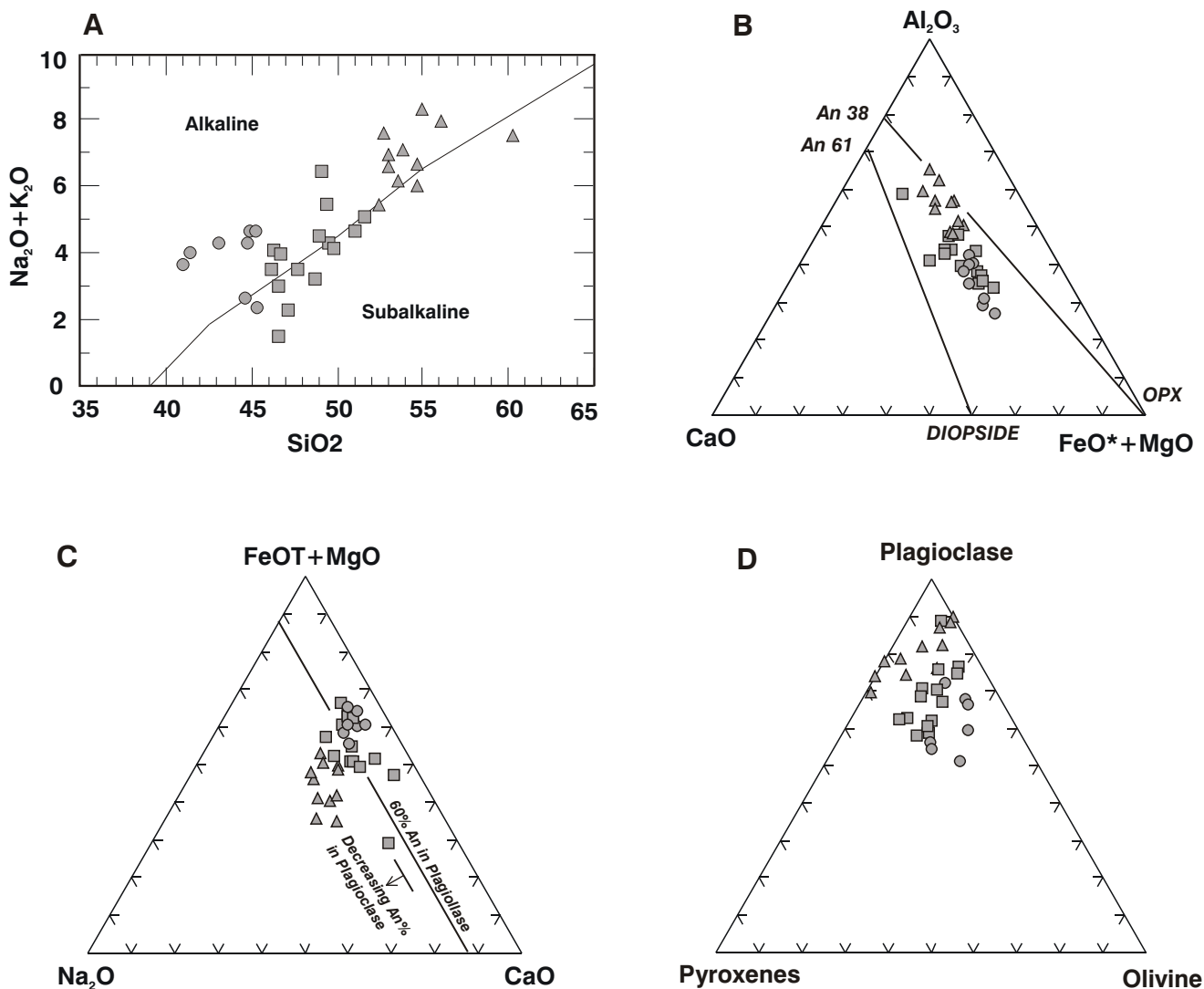


Figure 3. Major-element and normative composition diagrams of the MMIS samples. Legend: circle - ultramafic rock ($\text{SiO}_2 < 46$ wt.%); square - mafic rock (46-52 wt.% SiO_2); triangle - intermediate rock ($\text{SiO}_2 > 52$ wt.%). (A) Plot of total alkali vs silica showing the alkaline-subalkaline boundary of Irvine and Baragar (1971). (B) ACF diagram. (C) FNC diagram. (D) Ternary plot of normative plagioclase, pyroxenes (diopside + hypersthene), and olivine.

the geochemistry of the rocks cannot be directly discussed in terms of models applied to lava, considerably limiting the petrogenetic interpretation. Normative mineral mode (CIPW norm, Table 3) and composition diagrams are used to characterize the rocks and to show their similarities and differences.

MEALY MOUNTAINS INTRUSIVE SUITE

Mesoscopic examination of the MMIS gabbro and diorite samples indicates they are essentially composed of plagioclase, clinopyroxene, orthopyroxene, accessory hornblende and biotite, and up to a few percent sulphides and Fe-Ti oxides; olivine occurs only rarely. Dioritic and mon-

zodioritic samples are distinguished by ubiquitous K-feldspar and by greater abundances of biotite and hornblende at the expense of pyroxenes.

Analyzed MMIS samples (Tables 1 and 2) have SiO_2 contents ranging from 41.01 to 60.24 wt.% (average = 49.29 wt.%; Table 1). Three subgroups are distinguished in Figures 3 and 4; ultramafic ($\text{SiO}_2 < 46$ wt.%), mafic (SiO_2 46 to 52 wt.%), and intermediate ($\text{SiO}_2 > 52$ wt.%). The mafic and ultramafic subgroups have average CaO of 9.80 wt.% (6.67 to 13.68 wt.%), Na₂O of 2.92 wt.% (1.37 to 4.37 wt.%), K₂O of 1.02 wt.% (0.14 to 2.25 wt.%), MgO of 6.56 wt.% (2.03 to 13.87 wt.%), and total iron expressed as Fe₂O₃ of 12.47 wt.% (5.06-19.90 wt.%). The intermediate subgroup, which

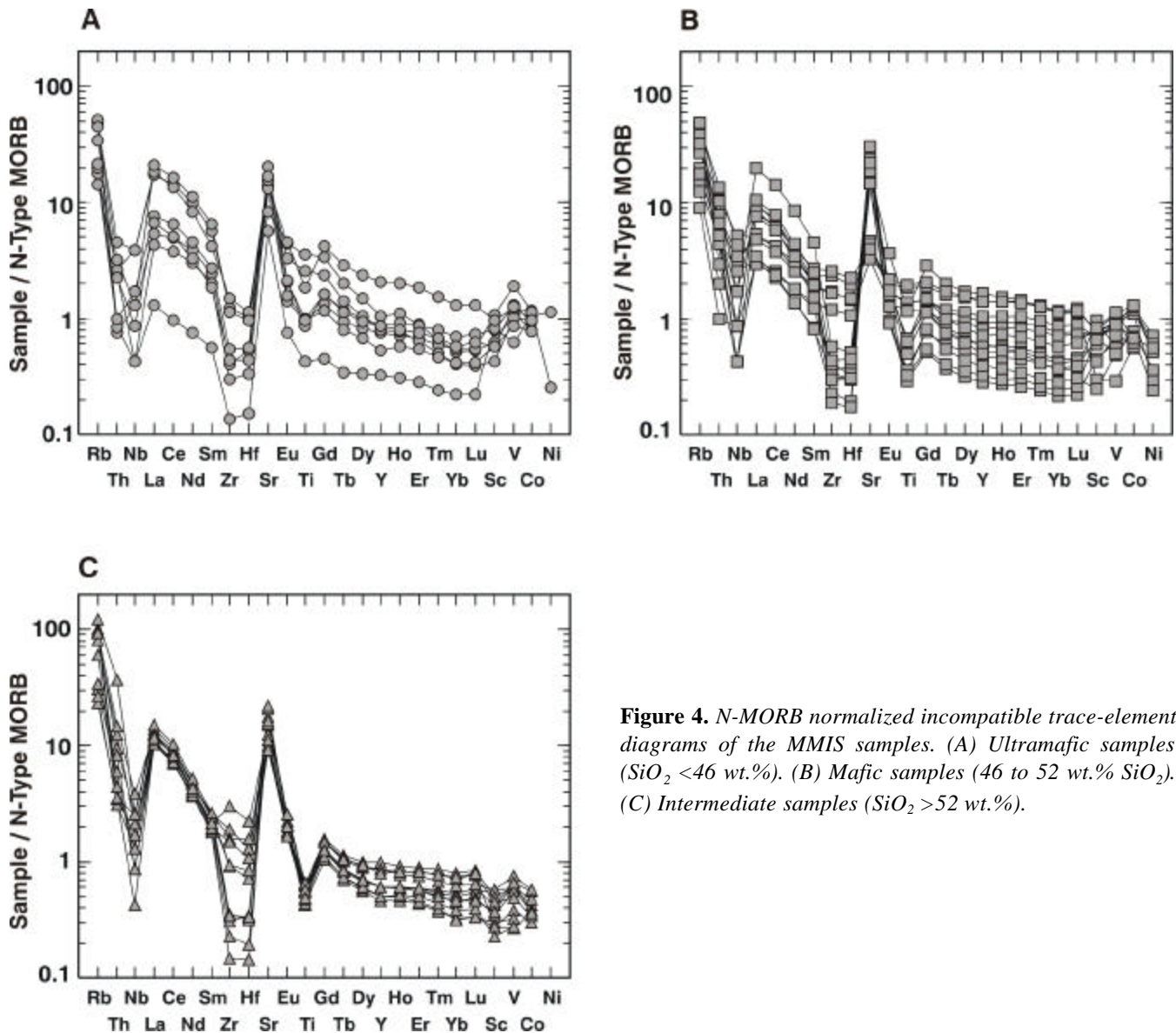


Figure 4. *N*-MORB normalized incompatible trace-element diagrams of the MMIS samples. (A) Ultramafic samples ($\text{SiO}_2 < 46$ wt.%). (B) Mafic samples (46 to 52 wt.% SiO_2). (C) Intermediate samples ($\text{SiO}_2 > 52$ wt.%).

has an average K_2O of 2.39 wt% (1.21 to 3.47 wt%), includes the monzodioritic samples. The latter are further distinguished by their higher average Na_2O of 4.55 wt% (4.06 to 5.04 wt%), and lower average CaO of 6.57 wt% (4.77 to 7.99 wt%), MgO of 3.09 wt% (1.83 to 4.90 wt%), and total iron as Fe_2O_3 of 7.19 wt% (5.30 to 9.39 wt%).

The mafic and ultramafic samples are also markedly enriched in Al_2O_3 (18.09 wt%) and deficient in TiO_2 (1.08 wt%) compared to the Mackenzie dyke swarm ($\text{Al}_2\text{O}_3 = 13.86$ wt%; $\text{TiO}_2 = 2.47$ wt%), to the Columbia River basalts ($\text{Al}_2\text{O}_3 = 14.34$ wt%; $\text{TiO}_2 = 2.33$ wt%), and to ocean-island basalts ($\text{Al}_2\text{O}_3 = 12.80$ wt%; $\text{TiO}_2 = 2.57$ wt%). This is consistent with crystal fractionation and accumulation.

All the analyzed MMIS samples are silica-saturated (olivine + hypersthene in the norm) or silica-undersaturated (olivine and nepheline in the norm), except Sample 5 containing over 60 wt.% SiO_2 (Table 3). They can be described as transitional to alkaline because they tend to plot near or slightly above the alkaline–subalkaline dividing line on an $\text{Na}_2\text{O} + \text{K}_2\text{O}$ vs SiO_2 diagram (Figure 3a). The abundance of normative olivine contrasts with its scarcity in the field; this results essentially from the anhydrous nature of the norm. The abundance of primary hornblende and biotite, which are common accessory minerals in these rocks, was likely controlled by the availability of deuterium during crystallization.

That the MMIS samples are essentially mixtures of plagioclase, pyroxenes and olivine is well illustrated in the ACF and normative Plagioclase–Pyroxenes–Olivine diagrams (Figures 3a and d). Normative plagioclase content of the ultramafic, mafic and intermediate subgroups average 49 wt.%, 59 wt.%, and 64 wt.% respectively. Orthose content of intermediate samples average 14 wt.%, in contrast with the 6 wt.% averages of the more mafic subgroups. Mean anorthite content in plagioclase decreases from approximately An60 in ultramafic samples, to approximately An56 and An40 in mafic and intermediate samples, respectively. Except for Sample 23, rocks of the mafic and ultramafic subgroups can be described either as olivine–gabbro or olivine–diorite (normative plagioclase < An50). Given their high normative orthose and K₂O content, intermediate samples are best described as olivine–monzodiorite. The trend of decreasing normative anorthite in plagioclase is reflected in the ACF and Na₂O–CaO–FeOT* +MgO diagrams (Figure 3b, and c), which also illustrate the limited compositional variability and the continuous nature of the suite.

The principal features of the trace-element patterns for the three subgroups are shown in N-MORB normalized extended elemental plots (Figure 4). The elemental abundances and distribution patterns of the three subgroups are remarkably similar, essentially overlapping each other. The pattern shows a marked enrichment in high-field strength elements (HFSE) relative to N-MORB, and a relatively flat heavy-REE distribution near or somewhat below the normalizing values. As expected, the samples of the intermediate subgroup are depleted in transition metals relative to the more mafic rocks. The close similarities observed in the trace-element pattern provide strong support to the hypothesis that these rocks are all genetically related.

The most notable features of the patterns are given by the well-defined negative Th, Nb, Zr, and Hf anomalies, which are coupled with a positive Sr spike. Titanium also defines a small negative anomaly, relative to Eu and Gd, that is more pronounced in the intermediate samples. The absence of an Eu anomaly, relative to Sm and Gd, is also notable. Such patterns indicate that fractional crystallization could have been an important process in the petrogenesis of the suite. The compatibility of Sr in plagioclase, which cotectically crystallized with olivine and/or pyroxenes, results in a nearly constant Sr in magma as elements highly incompatible in plagioclase, olivine, and pyroxenes gradually increase concentration in residual melt.

MECATINA TERRANE GABBRO (Unit M_{MC} gbr)

The geochemistry of gabbro bodies and dykes in the Mecatina terrane differs considerably from the geochem-

istry of MMIS gabbro and gabbro-norite. Samples of Mecatina terrane gabbro have a major-element composition that closely resembles the typical gabbro and basalt of Le Maitre (1976). The presence of plagioclase cumulate in the rock is suggested by the small enrichment in Al₂O₃ (16.37 wt.%) and depletions in TiO₂ (1.71 wt.%) and P₂O₅ (0.23 wt.%) compared to the Mackenzie dyke swarm (Al₂O₃ =13.86 wt.%; TiO₂= 2.47 wt.%; P₂O₅=0.97 wt.%), and to the Columbia River basalts (Al₂O₃ =14.34 wt.%; TiO₂= 2.33 wt.%; P₂O₅=0.45 wt.%).

The N-MORB normalized extended elemental plots also show a marked enrichment in high-field strength elements (HFSE), and very gently dipping heavy-REE (Figure 5a). In contrast to MMIS rocks, there is no depletion in Th, Zr, and Hf, and the amplitudes of the Nb, Sr, and Ti anomalies are much smaller. While the least fractionated samples exhibit prominent positive Sr spikes, and small positive Eu anomalies, the most fractionated rocks show much smaller positive Sr anomalies, and weak negative Eu and Ti anomalies. These signatures suggest that the samples with the strongest positive Eu anomalies and the lowest concentrations of incompatible elements are likely to represent plagioclase cumulate and, conversely, that those with higher concentrations of incompatible elements are most likely to approach residual liquid composition. Accordingly, the compositional range of samples having higher concentrations of incompatible elements is likely to reflect the composition of the parent magma.

It is emphasized that the patterns suggest limited Ti fractionation, and show no evidence for Th, Zr, Hf and Y fractionation, hence providing justification for the use of incompatible trace-element tectonic discrimination diagrams. The Mecatina terrane gabbro samples plot within the 10 percent probability ellipse for volcanic-arc basalts of Pearce (1996) in both the Ti–Zr–Y (Pearce and Cann, 1973) and Th–Ta–Hf (Wood *et al.*, 1979) discrimination diagrams (Figure 5b, and c). This is consistent with emplacement in an Andean-type arc setting.

QUARTZ MONZONITE (Unit M_{MC} qmm), PORPHYRITIC GRANITE (Unit M_{MC} kpg), and GRANITE (Unit M_{LG} kpg-grn)

Given the small number of analyses, interpretation of the geochemical data from the early Mesoproterozoic (Pinwarian) Mecatina terrane quartz monzonite (Unit M_{MC} qmm) and K-feldspar porphyritic granite (Unit M_{MC} kpg), and of the late- to post-Grenvillian granite (M_{LG} kpg-grn) samples is provisional. Significant differences in trace-element compositions are nonetheless recognizable (Figure 6a). Continental crust normalized REE patterns of the Mecatina terrane quartz monzonite and K-feldspar por-

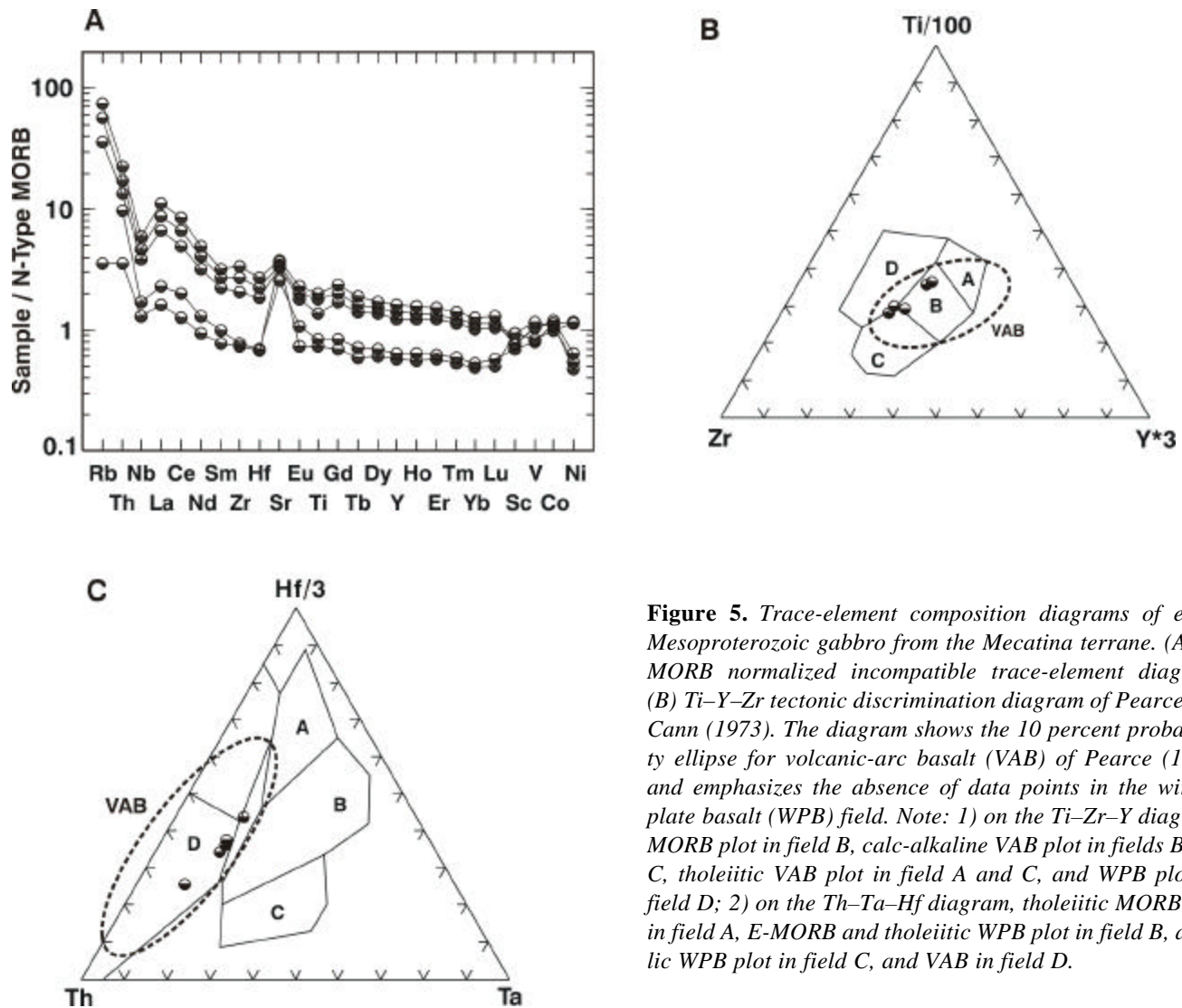


Figure 5. Trace-element composition diagrams of early Mesoproterozoic gabbro from the Mecatina terrane. (A) N-MORB normalized incompatible trace-element diagram. (B) Ti–Y–Zr tectonic discrimination diagram of Pearce and Cann (1973). The diagram shows the 10 percent probability ellipse for volcanic-arc basalt (VAB) of Pearce (1996) and emphasizes the absence of data points in the within-plate basalt (WPB) field. Note: 1) on the Ti–Zr–Y diagram, MORB plot in field B, calc-alkaline VAB plot in fields B and C, tholeiitic VAB plot in field A and C, and WPB plot on field D; 2) on the Th–Ta–Hf diagram, tholeiitic MORB plot in field A, E-MORB and tholeiitic WPB plot in field B, alkalic WPB plot in field C, and VAB in field D.

phyritic granite show subtle enrichments in relative light- to heavy-REE, with negative Eu anomalies varying in amplitude. The width of the trace-element spectrum reflects compositional variability of these units. In contrast, late- to post-Grenvillian granite samples are singled out by their generally higher abundances in trace elements, a small positive Ce anomaly, and an overall greater enrichment in relative light- to heavy-REE. It may be worth noting that granite samples from sites 38 and 53 share the geochemical signature of the early Mesoproterozoic granitic rocks, although initially assigned to other units from field observations (Figure 6a).

Analyses from the early Mesoproterozoic units plot within or close to the boundary of volcanic-arc granite fields on both Nb–Y and Rb–Y+Nb tectonic discrimination diagrams (Figure 6b and c). This is consistent with models

involving the development of an early Mesoproterozoic Andean-type margin along the southeastern margin of Laurentia. In contrast, the late- to post-Grenvillian granites form a tight cluster in the within-plate-granite field, consistent with their emplacement in significantly older crust.

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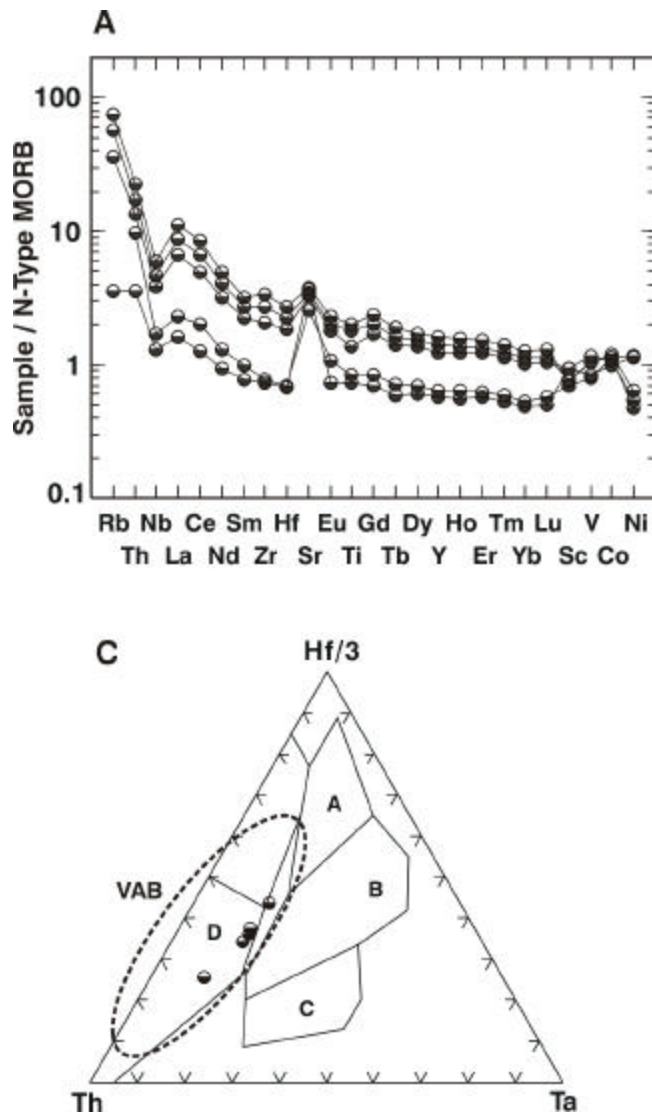


Figure 6. Trace-element composition diagrams for granitic samples. Legend: triangle - late- to post-Grenvillian K-feldspar porphyritic granite (M_{LG} kpg), closed square - quartz monzonite (M_{MC} qmm) and K-feldspar porphyritic granite (M_{MC} kpg) from Mecatina terrane, and open square - granite sample from Site 39 of uncertain affinity. (A) Continental crust normalized REE patterns. (B) Nb vs Y tectonic discrimination diagram of Pearce et al., (1984). (C) Rb vs Y+Nb tectonic discrimination diagram of Pearce et al., (1984). WPG - within-plate granite; ORG - ocean-ridge granite; VAG - volcanic-arc granite; Syn-COLG - syn-collision granite.

REFERENCES

Emslie, R.F.

1976: Mealy Mountains complex, Grenville Province, southern Labrador. *In* Report of Activities, Part A. Geological Survey of Canada, Paper 76-1A, pages 165-170.

Emslie, R.F. and Hunt, P.A.

1990: Ages and petrogenetic significance of igneous-charnockite suites associated with massif anorthosites, Grenville Province. *Journal of Geology*, Volume 98, pages 213-231.

Gower, C.F.

1996: The evolution of the Grenville Province in eastern Labrador. *In* Precambrian Crustal Evolution in the North Atlantic Region. Edited by T.S. Brewer. Geological Society Special Publication No. 112, pages 197-218.

1999: Geology of the Crooks Lake map region, Grenville Province, eastern Labrador. *In* Current Research. Newfoundland Department of Mines and Energy, Geological Survey, Report 99-1, pages 41-58.

- Gower, C.F., Heaman, L.M., Loveridge, W.D., Schärer, U. and Tucker, R.D.
1991: Grenvillian granitoid plutonism in the eastern Grenville Province, Canada. *Precambrian Research*, Volume 51, pages 315-336.
- Gower, C.F. and van Nostrand, T.
1996: Geology of the southeast Mealy Mountains region, Grenville Province, southeast Labrador. *In Current Research. Newfoundland Department of Mines and Energy, Geological Survey, Report 96-1*, pages 55-71.
- Irvine, T.N. and Baragar, W.R.A.
1971: A guide to the chemical classification of the common volcanic rocks; *Canadian Journal of Earth Sciences*, Volume 8, pages 523-548.
- James, D.T.
1999: Geology of the Kenamu River area (13C/NE), Grenville Province, southern Labrador. *Geological Survey of Newfoundland and Labrador, Department of Mines and Energy, Open File 013C/0040*, 1:100 000-scale map with descriptive notes.
- James, D.T., Kamo, S. and Krogh, T.E.
2000: Preliminary U–Pb geochronological data from the Mealy Mountains Terrane, Grenville Province, southern Labrador. *In Current Research. Newfoundland Department of Mines and Energy, Geological Survey, Report 2000-1*, pages 169-178.
- James, D.T., Kamo S., Krogh T. and Nadeau L.
This volume: Preliminary U–Pb geochronological data from Mesoproterozoic rocks, Grenville Province, southern Labrador.
- James, D.T. and Lawlor, B.
1999: Geology of the Grenville Province, Kenamu River area (NTS 13C/NE), southern Labrador: Preliminary observations of Labradorian and Pre-Labradorian(?) intrusions. *In Current Research. Newfoundland Department of Mines and Energy, Geological Survey, Report 99-1*, pages 59-70.
- James, D.T. and Nadeau, L.
2000a: Geology of the Minipi Lake area (NTS 13C/South): New data from the southern Mealy Mountains Terrane, Grenville Province, Labrador. *In Current Research. Newfoundland Department of Mines and Energy, Geological Survey, Report 2000-1*, pages 179-196.
- 2000b: Geology of the Minipi Lake area (NTS 13C/South), Grenville Province, southern Labrador. *Geological Survey of Newfoundland and Labrador, Department of Mines and Energy, Open File 013C/0041*, 1:100 000-scale map with descriptive notes.
- Nunn, G.A.G. and van Nostrand, T.
1996: Geology of the Kenemich River map area (NTS 13G/SW), Labrador. *In Current Research. Newfoundland Department of Mines and Energy, Geological Survey, Report 96-1*, pages 73-83.
- Pearce, J.A.
1996: A user's guide to basalt discrimination diagrams. *In Trace Element Geochemistry of Volcanic Rocks: Applications for Massive Sulphide Exploration. Edited by D.A. Wyman. Geological Association of Canada, short course notes, Volume 12*, pages 79-113.
- Pearce, J.A. and Cann, J.R.
1973: Tectonic setting of basic volcanic rocks determined using trace element analyses. *Earth and Planetary Science Letters*, Volume 19, pages 290-300.
- Wardle, R.J., Gower, C.F., Ryan, B., Nunn, G.A.G., James, D.T. and Kerr, A.
1977: Geological map of Labrador; 1:1 million scale. *Government of Newfoundland and Labrador, Department of Mines and Energy, Geological Survey, Map 97-07*.
- Wood, D.A., Joron, J.-L. and Treuil, M.
1979: A re-appraisal of the use of trace elements to classify and discriminate between magma series erupted in different tectonic settings. *Earth and Planetary Science Letters*, Volume 45, pages 326-336.